

# Parametric approach to promote a divergence-free flow in the image-based motion estimation with application to bioirrigation

Naratip Santitissadeekorn<sup>1</sup>, Christof Meile<sup>2</sup>, Erik Bollt<sup>3</sup>, and George Waldbusser<sup>4</sup>

<sup>1</sup>Department of Mathematics, University of Surrey, Guildford, UK

<sup>2</sup>Department of Marine Sciences, The University of Georgia, Athens, GA, USA

<sup>3</sup>Department of Mathematics and Computer Science, Clarkson University, Potsdam, NY, USA

<sup>4</sup>College of Earth, Ocean, and Atmospheric Sciences, Oregon State University, Corvallis, OR, USA

May 5, 2022

## Abstract

Flow fields are determined from image sequences obtained in an experiment in which benthic macrofauna, *Arenicola marina*, causes water flow and the images depict the distribution of a tracer that is carried with the flow. The experimental setup is such that flow is largely two-dimensional, with a localized region where the *Arenicola* resides, from which flow originates. Here, we propose a novel parametric framework that quantifies such flow that is dominant along the image plane. We adopt a Bayesian framework so that we can impart certain physical constraints on parameters into the estimation process via prior distribution. The primary aim is to approximate the mean of the posterior distribution to present the parameter estimate via Markov Chain Monte Carlo (MCMC). We demonstrate that the results obtained from the proposed method provide more realistic flows (in terms of divergence magnitude) than those computed from classical approaches such as the multi-resolution Horn-Schunk method. This highlights the usefulness of our approach if motion is largely constrained to the image plane with localized fluid sources.

## 1 Introduction

Determining motion from image sequences is a common inverse problem encountered in many natural sciences and engineering fields. Classical approaches such as the Horn-Schunk method [21] globally impose incompressibility to the entire image, since they are designed primarily for rigid-motion estimation in image processing or video analysis, and promote flow that varies smoothly

in space. A comprehensive survey of these methods is given in [3, 17]. Variational methods and algorithms for fluid motions were developed for the geophysical fluid motion estimation [1, 10, 12, 20, 39]. However, these methods are hardly a hand-in-glove fit for all general applications but rather designed with different emphases, e.g., vorticity estimates or temporal smoothness.

Optical flow approaches have also been used in oceanography, most commonly in the determination of surface flow from satellite imagery or derived products such as sea surface temperature [11, 27]. Here, we present an approach that is aimed at the quantification of fluid flow in marine sediments. Advective flow of porewater in marine sediments can be caused by a number of mechanisms [40], including (1) deeply sourced buoyancy causing the upward flow observed at hydrothermal vents or cold seeps (e.g. [4, 6]), (2) flow over topographic features, such as sand ripples [41] or (3) the activity of benthic macrofauna [50]. Termed bioirrigation, biologically induced flow can be a major influence on solute exchange between sediment and overlying water [30], organic matter cycling or sediment cracking [49]. Its impact on sediment biogeochemistry is evident from observations with planar optodes, revealing the profound, spatially and temporally dynamic imprint of the activity of macrofauna [34].

Benthic macrofauna are abundant in many marine sediments, with a typical mixing depth of about 10 cm [7, 44], and operate at scales that are amenable for ex situ laboratory studies. Bioirrigation is most commonly estimated from tracer concentration profiles, from which it is quantified as the process required to reconcile the observations with simulated concentration profiles after accounting for the contribution of diffusion, or known sources and sinks [16, 31]. More recently, fluid flow induced by bioirrigation has been estimated from images using (classical) correlation and gradient methods [14, 22]. Kaza [22] employed the classical Horn-Schunk [21] and Lucas-Kanade methods [24], which promote a globally or locally smoothly varying flow and do not provide a systematic way to measure reliability of a recovered velocity field.

In this study, we expand on that work and introduce a novel approach to quantify porewater flow induced by the activity of burrowing organisms in marine sediments. Our approach is tailored to the problem at hand in that it does not impose incompressibility to the entire image but applies a novel parametric approach and includes uncertainty quantification. We use a Bayesian approach similar to [42, 43], where the velocity field is probabilistically estimated by combining a model prediction simulated by a given fluid model with an actual observation (i.e. image sequences) [9, 28, 45]. It avoids the difficult task of deriving the Euler-Lagrange equations to achieve the desired result. The Bayesian framework also provides an uncertainty quantification of the estimated flow field that can

give an insight into the error statistics of the optical flow algorithm.

## 2 Data description and image preparation

A thin (ant-farm style) aquarium (0.22 m height  $\times$  0.445 m width  $\times$  0.022 m depth) was set up to study the effect of macrofauna on porewater flow. The aquarium was filled with approximately 0.16 m of near-shore muddy sand type sediment overlain by seawater. After letting the sediment settle for 1 week, a lugworm, *Arenicola marina*, an abundant macrofauna in the coastal ocean of Europe and North America, was placed in this aquarium. Lugworms are a few centimetres long, and are subsurface deposit-feeders who pump water through their burrows into the sediment (e.g. [35]), acting as ecosystem engineers [47]. Lugworm inject fluid into the sediment by inducing a peristaltic tail-to-head moving pumping wave[36]. This is done at the feeding pocket, located at the dead end of their J-shaped burrows. When the lugworms are pumping, water is moved through this structure and then injected into the sediment at the dead-end. In previous work, similar organisms (*Abarenicola marina*, a larger relative to *Abarenicola pacifica* with similar behavior, see e.g. [53]), have been modeled with radius of the feeding pocket of 2.5mm [33]. Given this small size, we describe this as a point source of water originating from outside the modeled domain consistent with observations and brought to the injection location via a burrow that is behind the experimentally observational plane. For the short time interval (on the order of minutes) covered by the images analyzed in this study, the burrow structure and injection location is considered constant. Qualitatively, flow velocities decrease with distance from the injection site and depend on the spatial variations in sediment permeability. As lugworms pump intermittently [54], the resulting flow velocities are temporally variable.

The lugworm was allowed to acclimate for 7 days. During this time water supply was maintained and air was bubbled in the overlaying water. The aquarium was maintained in the dark.

The flow induced by the lugworm’s activity was visualised by adding fluorescein, a dissolved fluorescent tracer, through a thin pipette to a location in the sediment close to where the burrowing animal resided. Its redistribution in the sediment - caused by pumping of the arenicolid worm - was captured as a series of high-resolution images of the fluorescent flow tracer. The aquaria was illuminated with blue led lights at an approximate wavelength of 450 nm, and a yellow barrier ‘photographic gel’ filter was used to reduce the blue light captured in the images. A 1 g per liter of fluorescein was added to a 32 ppt reagent grade NaCl solution to match the density of the resident pore water (fluid). RGB (red, green, blue) images were captured with a Fujifilm FinePix S5Pro

model at resolution  $2848 \times 4256$  pixels every 10 s. These images are subsequently referred to as *real RGB images*, and three out of the 41 images analyzed are displayed in Figure 1.

The use of a thin aquarium ensures that animal activity leads to a flow visible on the aquarium surface and the assumption of a two-dimensional optical flow. Thus, the incompressible 2D flow is expected to be largely divergence free, except near the location where the fluid is injected, which may not be parallel to the image plane.

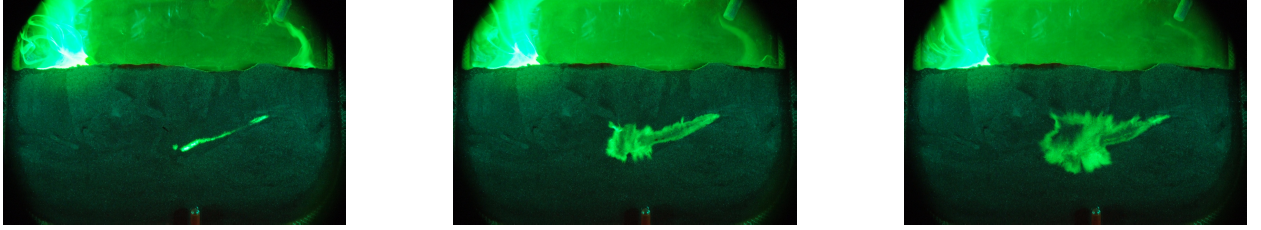


Figure 1: Real RGB images depicting spatial tracer distribution in a  $0.22 \text{ m} \times 0.445 \text{ m} \times 0.022 \text{ m}$  aquarium where bioirrigation is simulated at the relative times  $t_0 = 0 \text{ s}$ ,  $t_1 = 190 \text{ s}$  and  $t_2 = 390 \text{ s}$ .

The real RGB images are converted to 8 bit grey scale using the linear combination [51]

$$F(x, y, t) = 0.2989\mathcal{R}(x, y, t) + 0.5870\mathcal{G}(x, y, t) + 0.1140\mathcal{B}(x, y, t), \quad (2.1)$$

where  $\mathcal{R}(x, y, t)$ ,  $\mathcal{G}(x, y, t)$  and  $\mathcal{B}(x, y, t)$  denote red, green and blue colour channel intensities. The coefficients in (2.1) using the Luma rec 601 conversion implemented in the MATLAB *rgb2gray* command. Arrays arising from (2.1) are subsequently referred to as *real grey scale images*, and a subset of them are depicted in Figure 2. Importantly, the tracer concentration  $C$  is proportional to  $F$ .

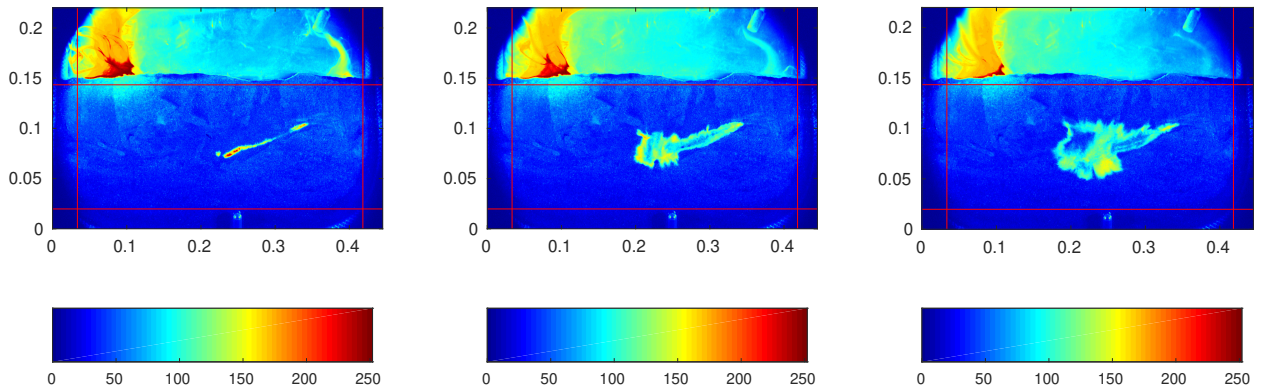


Figure 2: False color gray scale images, whose intensities are obtained using the linear combination in (2.1) with red, green and blue color channels depicted in Figure 1, and cropping boundaries overlaid on top in red.

As we aimed at the quantification of flow in the subsurface, our analyses are carried out on the area delineated in red in Figure 2 only, thus removing the overlying water (top 0.16 m) and a small area



near the bottom and sides of the aquarium, where the external lighting led to visible variations in brightness that might negatively impact optical flow computations. This image cropping reduces resolution to  $1600 \times 3680$  m, resulting in  $> 11$  million velocity components at the pixel level. To lower computational burden, the real grey scale images were down-sampled by a factor of eight using bilinear interpolation to a resolution of  $200 \times 460$ .

### 3 Mathematical Modeling

Consider the *Fokker-Planck equation* in (3.1), where  $C(x, y, t)$  denotes tracer concentration,  $\mathbf{u}(x, y, t)$  is the velocity vector, and  $\nabla = \begin{bmatrix} \frac{\partial}{\partial x} & \frac{\partial}{\partial y} \end{bmatrix}^T$  is the gradient operator [32, 33].

$$(\phi C)_t = \nabla \cdot (\phi D_e \nabla C - \phi \mathbf{u} C). \quad (3.1)$$

Let  $\phi(x, y, t)$  denote sediment *porosity*, the ratio of void to total volume, on the order of 0.7 in sandy sediments inhabited by *Arenicola marina*. Here, it is assumed to be a spatio-temporally constant (for simplicity), so that

$$C_t = \nabla \cdot (D_e \nabla C - \mathbf{u} C). \quad (3.2)$$

The scalar  $D_e \in \mathbb{R}$  in (3.2) denotes the *effective diffusion* of the sediment, and is given by

$$D_e = \frac{D}{1 - 2 \ln \phi}, \quad (3.3)$$

where  $D \in \mathbb{R}$  is the *molecular diffusion coefficient* in aqueous solution. For fluorescein,  $D = 4.25 \times 10^{-2} \text{ m}^2 \text{ s}^{-1}$  so that  $D_e = 2.48 \times 10^{-10} \text{ m}^2 \text{ s}^{-1}$  [13]. Consider the *Péclet number* given by

$$\text{Pe} = \frac{\ell \|\mathbf{u}\|_2}{D_e}, \quad (3.4)$$

where  $\ell > 0$  is a length scale and  $\|\cdot\|_2$  denotes a Euclidean norm. Advection is the more prominent solute transport mechanism when  $\text{Pe} \gg 1$ . Now  $\ell$  is  $\mathcal{O}(10^{-2})$  m, in line with the dimensions of the aquarium from which the real RGB images arise. Also, previous research suggests that the flow speed  $\|\mathbf{u}\|_2$  of water induced by bioirrigating *Arenicola marina* is approximately  $\mathcal{O}(10^{-6}) \text{ m s}^{-1}$ – $\mathcal{O}(10^{-5}) \text{ m s}^{-1}$  [22, 48]. This, alongside the earlier definitions of porosity and diffusion coefficient, indicates that  $\text{Pe} \gg 1$ . This motivates the assumption that diffusion is negligible, particularly in

the region of interest in the vicinity of the burrow, so that (3.2) can be rewritten as

$$-C_t = \nabla \cdot (\mathbf{u}C). \quad (3.5)$$

We will assume incompressibility and the condition  $\nabla \cdot \mathbf{u} = 0$ , whose validity for the current context was discussed in previous research [22, 48]. It then follows that

$$\begin{aligned} -C_t &= C\nabla \cdot \mathbf{u} + \mathbf{u} \cdot \nabla C. \\ &= \mathbf{u} \cdot \nabla C. \end{aligned} \quad (3.6)$$

The connection between (4.1) and (3.6) are made by assuming that tracer concentration is proportional to grey scale intensity (i.e.  $C \propto F$ ) [22].

## 4 Horn-Schunck Method

For a velocity field to be estimated, a mathematical expression relating it to grey scale intensity is required. Consider a pair of images, whose grey scale intensities are defined by  $F_0(x, y, t_0)$  and  $F_1(x, y, t_1)$  for  $t_0 < t_1$ . An *optical flow*  $\mathbf{u}(x, y, t) = [u(x, y, t) \ v(x, y, t)]^T$  is an apparent motion field mapping one image to the next. Conventional optical flow problem relies on the (pixel) brightness constancy approximation. This is expressed mathematically in (4.1), where the subscripts  $x$ ,  $y$  and  $t$  denote derivatives with respect to the  $x$ -axis,  $y$ -axis and time [21].

$$F_t + uF_x + vF_y = 0. \quad (4.1)$$

In the current context, this approximation can be justified by requiring the null-divergence of the velocity in (3.5) at every image pixel. This is, however, ill-posed since two velocity components have to be determined from one equation. To address this issue, a classical *Horn-Schunck method* was developed based on a variational principle [21]. In particular, the Horn-Schunck method further constrains the optical flow to be spatially smooth, which can be expressed mathematically as

$$\|\nabla \mathbf{u}\|_2^2 + \|\nabla \mathbf{v}\|_2^2 < \infty. \quad (4.2)$$

Locating the optical flow that minimises (4.1) subject to (4.2), with respect to the spatial region  $\Omega$  that the images depict, can be expressed mathematically as

$$\mathbf{u} = \underset{\mathbf{u}}{\operatorname{argmin}} \int_{\Omega} \underbrace{(\mathbf{u} \cdot \nabla \mathbf{F} + \mathbf{F}_t)^2 + \lambda^2 (\|\nabla \mathbf{u}\|_2^2 + \|\nabla \mathbf{v}\|_2^2)}_{\mathcal{I}} d\Omega \quad (4.3)$$

where  $\lambda \in \mathbb{R}$  is a *smoothing parameter*. Following [21], the Euler- Lagrange equations for (4.3) are given by the following coupled system of PDEs, where  $\Delta$  is the Laplacian operator,

$$\begin{aligned} F_x(F_x u + F_y v + F_t) &= \lambda^2 \Delta u, \\ F_y(F_x u + F_y v + F_t) &= \lambda^2 \Delta v. \end{aligned} \quad (4.4)$$

At the image pixel level, the system in (4.4) can be numerically solved in which the Laplacian is approximated by

$$\Delta u \approx \bar{\bar{u}} - u, \quad \bar{\bar{u}} = \frac{1}{12} \begin{bmatrix} 1 & 2 & 1 \\ 2 & 0 & 2 \\ 1 & 2 & 1 \end{bmatrix} * u. \quad (4.5)$$

The  $*$  in (4.5) means a 2D convolution on the image pixel. In other words,  $\bar{\bar{u}}$  is a local weighted average. After substituting the Laplacian approximation in (4.5) into (4.4), we can apply the Jacobi iteration to obtain the following iterative scheme, where the subscript  $i$  denotes iteration indices,

$$\begin{aligned} u_{i+1} &= \bar{\bar{u}}_i - \frac{F_x(F_x \bar{\bar{u}}_i + F_y \bar{\bar{v}}_i + F_t)}{\lambda^2 + F_x^2 + F_y^2}, \\ v_{i+1} &= \bar{\bar{v}}_i - \frac{F_y(F_x \bar{\bar{u}}_i + F_y \bar{\bar{v}}_i + F_t)}{\lambda^2 + F_x^2 + F_y^2}. \end{aligned} \quad (4.6)$$

The differential nature of the Horn-Schunk method can make it difficult to estimate a large displacement and not robust to noise. In practice, a multi-resolution frame work (e.g. see [5, 37]) is commonly used to mitigate these issues. This approach is based on an incremental estimate that refines a previous estimate using finer resolution image at the next higher resolution level, which is commonly referred to as “pyramid” level. At the top of the pyramid level, the estimation is carried out at the original resolution. At the lower levels, the image is repeatedly down-sampled for a pre-specified factor. The details of the multi-resolution Horn-Schunk method is discussed in Appendix A. The number of pyramid levels is chosen as  $\mathcal{L} = 3$  here to obtain the results below. Ideally, magnitudes of additional terms added to a current correction increment in (A.3) should depend on how faithfully the associated optical flow mimics the unknown true solution. So if a

current correction increment produces an optical flow that almost exactly mimics the true solution, then magnitudes of additional terms in (A.3) at subsequent iterations should be approximately zero. The magnitudes of the quotients in (A.3) can be controlled by the value of  $\lambda$ .

The smoothing parameter should not be so “small” that these quotients are considered “big” in regions where contrast is low. Otherwise, optical flow speeds could then temporally increase to a point where they are viewed as being unrealistically “large”. However, the smoothing parameter should not be too “large”. This could induce an unrealistic high level of spatial smoothing. By trial and error, the value of  $\lambda = 50$  is chosen because it appears to be one of the “smaller” smoothing parameter values where flow speeds do not become unrealistically “large” over time. This is examined by plotting a frequency distribution of flow speeds at each relative time point in Figure 3.

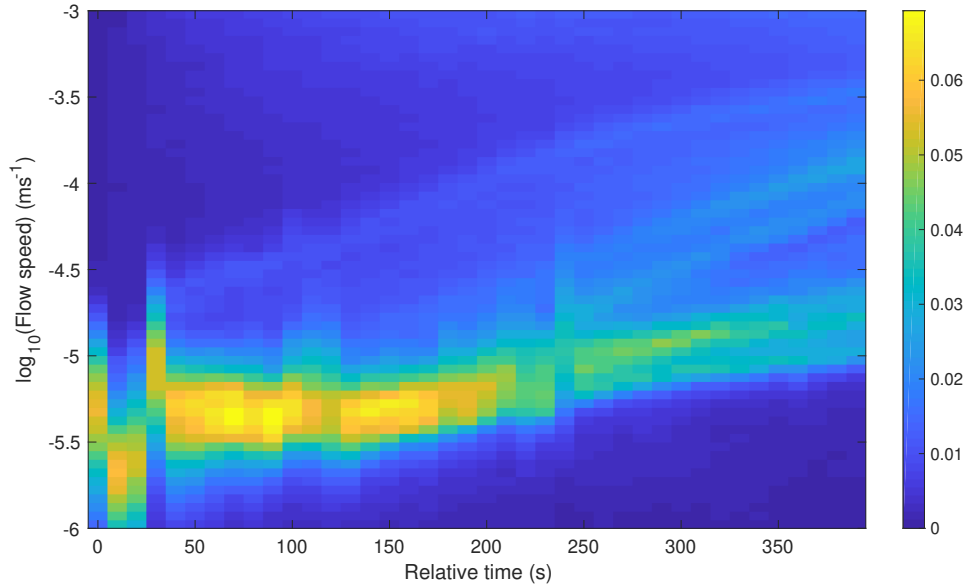


Figure 3: Each column shows the frequency distribution of optical flow speeds obtained by the multi-resolution Horn-Schunck method for  $\lambda = 50$  and  $M = 1$ .

The number of iterations of (A.3) at each pyramid level, denoted by  $M$ , also needs to be defined. The results here are shown for  $M = 1$  as per the following justification. It appears that when using  $M > 1$ , at later relative time points, the only region where flow points downwards is between the likely location of the organism and the bottom edge of the down-sampled and cropped real grey scale images. This suggests that there is outflow and no inflow at the top boundary of these images, which roughly coincides with the sediment-water interface according to Figure 2. However, given that the sides and bottom of the aquarium are impermeable, net flow across the sediment-water interface should be zero. We expect some outflow in the regions near the aquarium walls were cropped due to uneven lighting (Section 2). However, the qualitative assessment of tracer

distribution in the overlying water (Figure 2), indicates that this flow is limited. Thus,  $M$  is chosen to enforce a little inflow at later time points.

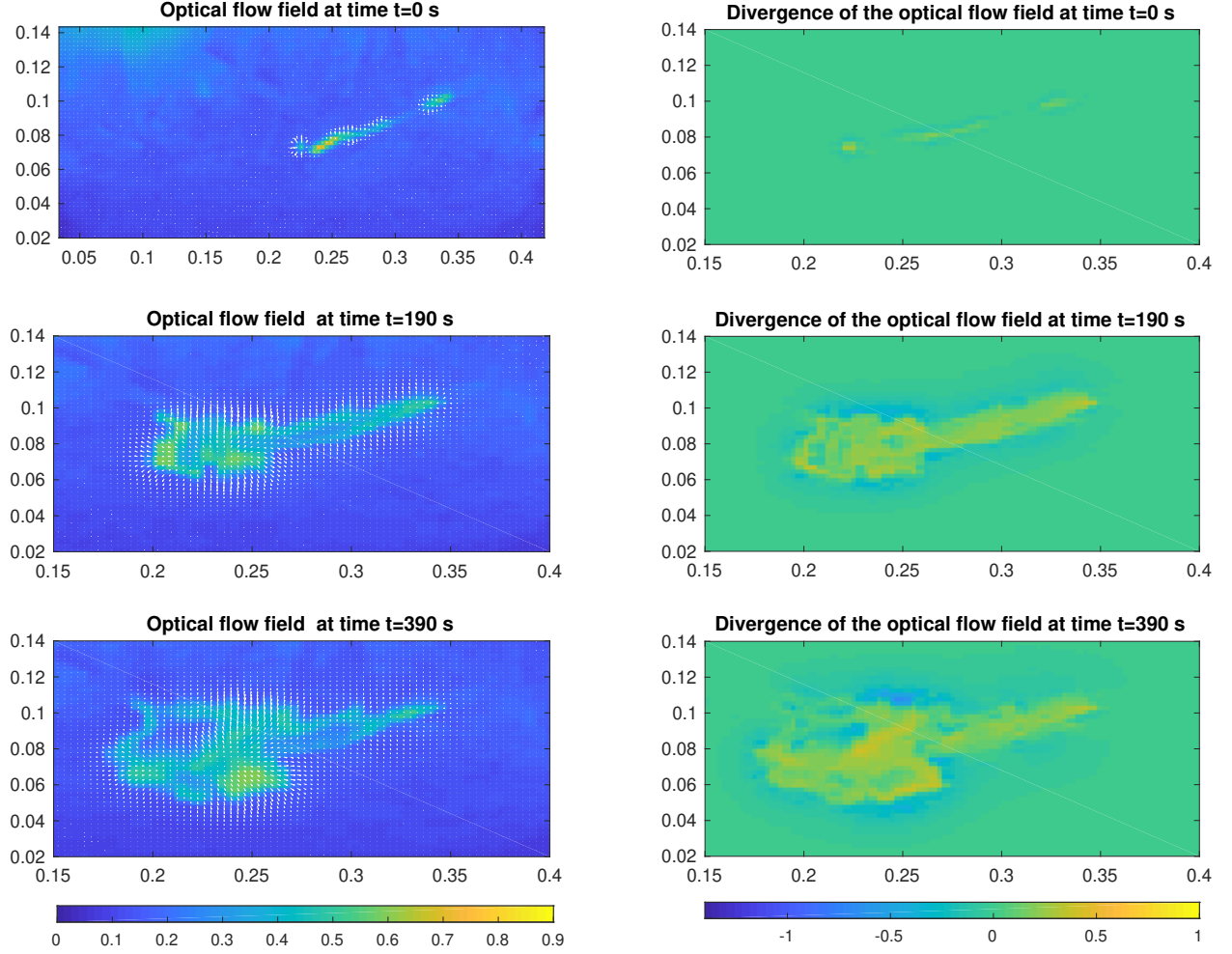


Figure 4: The first column depicts optical flows, denoted by white arrows, arising from the multi-resolution Horn-Schunck method when  $\lambda = 50$  and  $M = 1$  at the relative times  $t_0 = 0$  s (first row),  $t_1 = 190$  s (second row) and  $t_2 = 390$  s (third row). They are overlaid on top of the down-sampled and cropped real grey scale images associated with those relative times. The second column depicts the associated divergences. Note that all parts of this figure have been further down-sampled by a factor of four, for ease of presentation.

Figure 4 reveals undesirably “large” divergence magnitudes where they ought to be comparatively “small”, i.e., away from the location where the worm activity induces flow (see spreading arrows near the center of the top left panel). Furthermore, largest velocities are computed away from this point, with no clear source of fluid flow being discernible. Thus, additional ways of imposing that divergence magnitudes immediately tend to zero upon departure from the point of injection are investigated. This is developed in the next section.

## 5 Derivation of parametric framework

*Arenicola marina* inject water into the sediment at the dead end of often J-shaped burrows. Simulations of bioirrigation by the lugworm have shown that flow fields are largely unaffected by the burrow structure itself [32]. Hence, we consider flow in the sediment being driven by a fluid source located at a small injection pocket. Away from injection pocket, flow is considered two-dimensional, parallel to the imaging surface, due to closely spaced front and back wall of the thin aquarium. Therefore, the condition  $\nabla \cdot \mathbf{u} = 0$  is incorporated within the Bayesian framework through the prior distribution with the exception to this at the point  $(\mathbf{a}, \mathbf{b})$ , where the organism injects fluid. In particular, we will write an ansatz for the velocity potential  $\Phi(x, y, t)$  of a two-dimensional incompressible point source with strength  $c > 0$  centred on the point  $(\mathbf{a}, \mathbf{b}) \in \mathbb{R}^2$ , subject to boundary conditions as discussed in Appendix A. The injection point induces fluid expansion with strength  $c$ , suggesting that  $\nabla \cdot \mathbf{u} = c$ . All above requirements suggest the condition

$$\nabla \cdot \mathbf{u} = c\delta(x - \mathbf{a}, y - \mathbf{b}). \quad (5.1)$$

We can write (5.1) in terms of a velocity potential by using  $\mathbf{u} = \nabla\Phi$ .

$$\Delta\Phi = c\delta(x - \mathbf{a}, y - \mathbf{b}). \quad (5.2)$$

To deal with the delta function on the right-hand side of (5.2), we follow the method demonstrated in [2] to approximate the solution of (5.2) with that of

$$\Delta\Phi = \nu\Phi, \quad (5.3)$$

where  $\nu \in \mathbb{R}$  is an eigenvalue. The ansatz of velocity potential of a two-dimensional incompressible point source with strength  $c$  centred on the location  $(\mathbf{a}, \mathbf{b})$  is given by

$$\Phi(x, y) = \frac{c}{2\pi} \ln \sqrt{((x - \mathbf{a})^2 + (y - \mathbf{b})^2)}. \quad (5.4)$$



We use the separation of variables to solve (5.4), see Appendix B for the outline of the derivation,

$$\begin{aligned} \Phi(x, y) = & - \sum_{j=0}^{\infty} \sum_{k=0}^{\infty} \frac{4c}{L_x L_y \left[ \left( \frac{k\pi}{L_x} \right)^2 + \left( \frac{(1+2j)\pi}{2L_y} \right)^2 \right]} \cos \left( \frac{k\pi}{L_x} x \right) \cos \left( \frac{(1+2j)\pi}{2L_y} y \right) \\ & \times \cos \left( \frac{k\pi}{L_x} x \right) \cos \left( \frac{(1+2j)\pi}{2L_y} y \right), \end{aligned} \quad (5.5)$$

where  $[0, L_x] \times [0, L_y]$  is the area under computation, see Figure 12. The infinite sums in (5.5) is approximated with the finite sums with the first  $M_x$  terms for the inner sum and  $M_y$  for the outer sum. In what follows, we choose  $M_x = M_y = 200$ . Utilising the ansatz in (5.5) means that once the parameters  $\{a, b, c\}$  are estimated, flow at each grid cell can be obtained using the relation  $\mathbf{u} = \nabla \Phi$ .

## 6 MCMC for flow field estimation

The relation  $\mathbf{u} = \nabla \Phi$  can be used to rewrite (4.1) in terms of velocity potentials [1, 21, 25].

$$-F_t = F_x \Phi_x + F_y \Phi_y. \quad (6.1)$$

Substituting the finite approximation of (5.5) into (6.1) yields

$$\begin{aligned} -F_t \approx & \sum_{j=0}^{M_y} \sum_{k=0}^{M_x} \frac{4c}{L_x L_y} \cos \left( \frac{k\pi}{L_x} x \right) \cos \left( \frac{(1+2j)\pi}{2L_y} y \right) \left[ - \frac{F_x k\pi}{L_x} \sin \left( \frac{k\pi}{L_x} x \right) \cos \left( \frac{(1+2j)\pi}{2L_y} y \right) \right. \\ & \left. - \frac{F_y (1+2j)\pi}{2L_y} \cos \left( \frac{k\pi}{L_x} x \right) \sin \left( \frac{(1+2j)\pi}{2L_y} y \right) \right]. \end{aligned} \quad (6.2)$$

This acts as an observation system relating the parameters  $\{a, b, c\}$  contained within the state vector to observed negated temporal grey scale intensity derivatives. Let  $\mathbf{z} := -F_t$  and  $\mathbf{x} := [a, b, c]^T$  and abbreviate (6.2) by

$$\mathbf{z} = \mathcal{H}(\mathbf{x}). \quad (6.3)$$

We will investigate the uncertainty of the parameters for the three observations of  $\mathbf{z} := -F_t$  approximated from the (grey-scale) image pairs at the relative time point  $(t = 0, t = 10)$ ,  $(t = 10, t = 20), \dots, (t = 40, t = 50)$  s. These will be denoted by  $\mathbf{z}(0), \mathbf{z}(10), \dots, \mathbf{z}(40)$ , respectively. We will simulate the observation one at a time to sequentially sampling  $p(\mathbf{x}|\mathbf{z}(0)), p(\mathbf{x}|\mathbf{z}(0), \mathbf{z}(10)), \dots, p(\mathbf{x}|\mathbf{z}(0), \dots, \mathbf{z}(40))$

## 6.1 Setting likelihood function and prior distribution

To adopt a Bayesian framework, we need to assume a conditional distribution of  $\mathbf{z}$  given  $\mathbf{x}$  and the prior distribution of  $\mathbf{x}$ . For simplicity, we assume the likelihood function of the observation  $\mathbf{z}$  at each time point to have the kernel (i.e. the non-normalized part of the distribution)

$$p(\mathbf{z}|\mathbf{x}) \propto \exp\left(-\frac{1}{2}(\mathbf{z} - \mathcal{H}(\mathbf{x}))^\top \mathbf{R}^{-1}(\mathbf{z} - \mathcal{H}(\mathbf{x}))\right), \quad (6.4)$$

where  $\mathbf{R} = \sigma_2^2 \mathbf{I}$  and  $\mathbf{I}$  is the identity matrix of an appropriate size. An empirical value of  $\sigma_2^2 = 1 \times 10^{-3} \text{ s}^{-2}$  is adopted here following a numerical experiment on synthetic data with similar features to our current real-world data [29]. We use the same likelihood function for  $\mathbf{z}$  at all time points. The state vector now contains source strengths, and  $x$ - and  $y$ -coordinates, which have different physical properties. For the source  $x$ -coordinates, the horizontal distance from the source to the right-hand boundary  $\tilde{\mathbf{a}} \in \mathbb{R}$  is to be estimated instead of  $\mathbf{a}$ , where  $\tilde{\mathbf{a}} = \mathbf{L}_x - \mathbf{a}$ . We treat these parameters as being probabilistically independent, i.e.,

$$p(\mathbf{x}) = p(\tilde{\mathbf{a}})p(\mathbf{b})p(\mathbf{c}). \quad (6.5)$$

The assumption of independence is justified because the organism's location in the sediment should not affect the flow it induces, nor does the burrowing depth impact the lateral positioning.

We now describe the choice of the prior distribution at the first time point and will subsequently discuss how the prior for later time points are constructed. The marginal prior distribution for the source strength is assumed to be a Gamma distribution to ensure the positivity. The prior assumption  $E[\mathbf{c}] = 2 \times 10^{-5} \text{ m}^2 \text{ s}^{-1}$  is made because flow speeds residing approximately within the interval  $\mathcal{O}(10^{-6}) \text{ m s}^{-1} - \mathcal{O}(10^{-5}) \text{ m s}^{-1}$ , which appear to be induced by bioirrigating *Arenicola marina* in previous research [22, 48]. The marginal prior variance is taken to be  $\text{Var}[\mathbf{c}] = 2.5 \times 10^{-11} \text{ m}^4 \text{ s}^{-2}$ . It should be “large” enough that it can be reduced when observations are assimilated, whilst not being so “large” that the information quantified by the marginal prior mean is overpowered by observations. The marginal prior means are derived from the  $x$ - and  $y$ -coordinates of the pixel in the first down-sampled and cropped real grey scale image with the greatest intensity, because intuitively this seems to be a realistic location of fluid injection. Let  $\delta_x = \mathbf{L}_x^{-1}$  and  $\delta_y = \mathbf{L}_y^{-1}$ . The marginal prior distributions at the first relative time point for the parameters  $\mathbf{b}$  is assumed to be a Gaussian distribution with mean  $E[\mathbf{b}] = 0.07 \text{ m}$  and variance  $\text{Var}[\mathbf{b}] = \frac{5\delta_y^2}{2} \text{ m}^2$ . Finally, consider the parameter  $\tilde{\mathbf{a}}$ . A negative value of  $\tilde{\mathbf{a}}$  is unrealistic. As such, a Gamma marginal prior distribution

is employed for  $\tilde{\mathbf{a}}$ . So we assume a Gamma distribution with mean  $E[\tilde{\mathbf{a}}] = 0.2058$  m and variance  $9\delta x^2$  m<sup>2</sup>.

Note that the prior mean is set to be “small”, with a view to quantifying the fact that the source lies very close to the right-hand boundary. Their related standard deviations of  $p(\tilde{\mathbf{a}})$  and  $p(\mathbf{b})$  are “small” compared with the width and height of the aquarium to ensure that the source x-and y-coordinates recovered from the Metropolis-Hastings sampler remain “close” to the regions with the greatest intensity. With above prior assumption, the Metropolis-Hastings sampler can be used to sample the posterior distribution

$$p(\mathbf{x}|\mathbf{z}(\mathbf{0})) \propto p(\mathbf{z}(\mathbf{0})|\mathbf{x})p(\tilde{\mathbf{a}})p(\mathbf{b})p(\mathbf{c}). \quad (6.6)$$

The posterior distribution of the parameter vector  $\mathbf{x}$  given  $\mathbf{z}(\mathbf{0})$  could then be used as the prior density for the next observation  $\mathbf{z}(\mathbf{10})$ . However, we only have a sample of the posterior distribution, not the analytical form. To deal with this issue, we assume the independent prior for  $\mathbf{x}$  when assimilating  $\mathbf{z}(\mathbf{0})$  and then construct the prior distribution as above by using the posterior (marginal) mean and variance for each parameter  $\mathbf{a}, \mathbf{b}, \mathbf{c}$  given  $\mathbf{z}(\mathbf{0})$  instead. The process is then repeated for the rest of the observations.

## 6.2 Metropolis Hasting

We use the Metropolis-Hastings (MH) sampler [19] to sample (6.6). The MH algorithm involves accepting or rejecting a proposed sample with the acceptance probability

$$\gamma = \min \left\{ 1, \frac{p(\mathbf{x}^*|\mathbf{z})}{p(\mathbf{x}_i|\mathbf{z})} \frac{p(\mathbf{x}_i|\mathbf{x}^*)}{p(\mathbf{x}^*|\mathbf{x}_i)} \right\}, \quad (6.7)$$

where  $\mathbf{x}^*$  is a (new) proposed sample and  $p(\mathbf{x}^*|\mathbf{x}_i)$  is a proposal distribution of  $\mathbf{x}^*$  given the current sample  $\mathbf{x}_i$ . We assume the proposal distribution of the following form

$$\begin{aligned} p(\mathbf{x}^*|\mathbf{x}_i) &= p(\tilde{\mathbf{a}}^*, \mathbf{b}^*, \mathbf{c}^*|\tilde{\mathbf{a}}_i, \mathbf{b}_i, \mathbf{c}_i) \\ &= p(\tilde{\mathbf{a}}^*|\tilde{\mathbf{a}}_i)p(\mathbf{b}^*|\mathbf{b}_i)p(\mathbf{c}^*|\mathbf{c}_i). \end{aligned} \quad (6.8)$$

The marginal proposal distribution  $p(\tilde{\mathbf{a}}^*|\tilde{\mathbf{a}}_i)$  is chosen to be the Gamma distribution with mean  $\tilde{a}_i$  and variance is obtained by pre-multiplying marginal prior variance of  $\tilde{a}$  by a factor of 0.36 here. The marginal proposal distribution  $p(\tilde{\mathbf{c}}^*|\tilde{\mathbf{c}}_i)$  is also defined by the Gamma distribution with mean  $\tilde{c}_i$  and variance is obtained as above. A Gaussian marginal proposal distribution is defined for source

y-coordinates to enforce that injection can realistically take place either above or below the current chain location. In particular,  $p(\mathbf{b}^*|\mathbf{b}_i)$  is a Gaussian distribution with mean  $b_i$  and variance is the marginal prior variance of  $b$  multiplied by 0.36. The ad-hoc choice of the above pre-multiplying factor is chosen from a trial and error to strike a balance between the acceptance and convergence rates; when the proposal variance is too “small”, only “small” steps can be taken across the state space and convergence is slow. However, we make no attempt to optimally tune this factor here.

### 6.3 Convergence Diagnosis

In what follows we will run multiple Markov chains, starting in different regions of the state space, and consider both inter-chain mixing and intra-chain mixing for convergence diagnosis. Suppose that there are  $\mathcal{C} \in \mathbb{N}$  Markov chains. Each chain ends up of length  $N_{\text{spin}} + N$ , where the first  $N_{\text{spin}}$  particles constitute an initial transient called a *spin-up period*. The latter  $N$  particles constitute the ensemble that the Metropolis-Hastings sampler generates. The variance of the  $j$ -th Markov chain (for  $j \in \{1, \dots, \mathcal{C}\}$ ) is given by

$$s_j^2 = \frac{1}{\tilde{N} - 1} \sum_{i=1}^{\tilde{N}} (c_{ij} - \bar{c}_j)^2 \quad (6.9)$$

after the  $\tilde{N}$ -th particle is obtained. The terms  $c_{ij}$  and  $\bar{c}_j$  denote the  $i$ -th member of the  $j$ -th Markov chain (for  $i \in \{1, \dots, \tilde{N}\}$ ), and the mean of the  $j$ -th Markov chain (respectively). Note that  $s_j^2$  is written in terms of source strengths but it can be computed for source x-and y-coordinates as well. The expression in (6.9) is used to define the *within-chains variance*, which is given by [18]

$$\mathcal{W} = \frac{1}{\mathcal{C}} \sum_{j=1}^{\mathcal{C}} s_j^2. \quad (6.10)$$

The *between-chains variance*, denoted by  $\mathcal{B}$ , is defined by [18, 23]

$$\frac{\mathcal{B}}{\tilde{N}} = \frac{1}{\mathcal{C} - 1} \sum_{j=1}^{\mathcal{C}} (\bar{c}_j - \bar{c})^2. \quad (6.11)$$

The term  $\bar{c}$  in (6.11) denotes the overall mean across all  $\mathcal{C}$  Markov chains of source strengths. As suggested in [18, 23], the following ratio

$$\hat{R} = \sqrt{\frac{\text{Var}[\hat{\mathcal{C}}|\mathbf{z}]}{\mathcal{W}}} \quad (6.12)$$

could be used to judge when a set of Markov chains converges, where

$$\text{Var}[\hat{\mathbf{c}}|\mathbf{z}] = \left(1 - \frac{1}{\hat{N}}\right) \mathcal{W} + \frac{\mathcal{B}}{\hat{N}}. \quad (6.13)$$

Typically, it is suggested that the MCMC should be run until  $\hat{R} < 1.1$  [23].

## 6.4 Results: MCMC

To avoid cluttering the presentation of the result, results are shown only at the relative times  $t_0 = 0$  s,  $t_1 = 20$  s and  $t_2 = 40$  s. This experiment runs 6 Markov chains and the initial positions of these chains are independently drawn from the prior distribution explained in Section 6.1. We decide on using 6 chains after the fact that we obtained an acceptable value of  $\hat{R}$  in this experiment. The spin-up period consists of the first 500 particles and the overall length of each chain, including the spin-up period, is 2000.

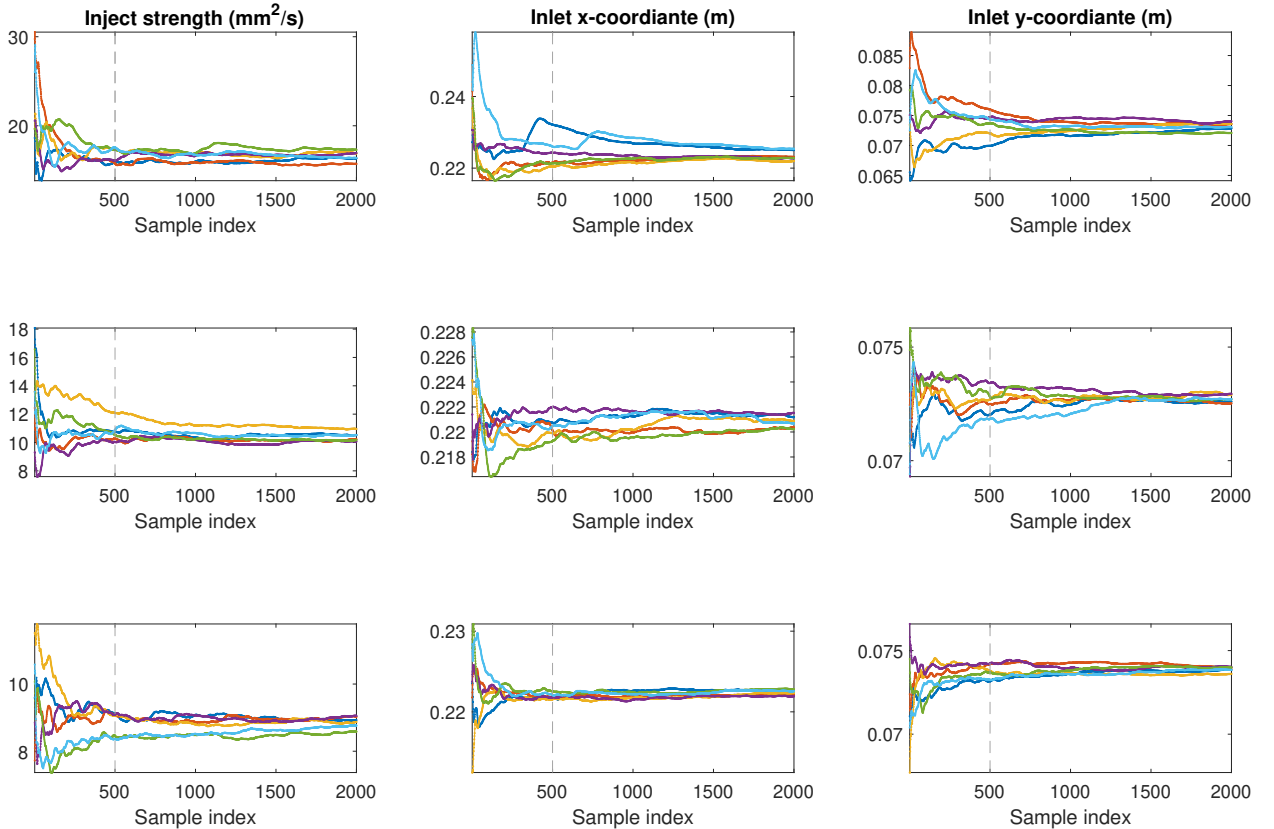


Figure 5: Evolving Markov chain means for source strengths (first column), x-coordinates (second column) and y-coordinates (third column). The relative times  $t_0 = 0$  s,  $t_1 = 20$  s and  $t_2 = 40$  s constitute the first, second and third rows (respectively). The dashed lines denote the end of the spin-up period, which constitutes the first 500 particles.

Figure 5 depicts evolving Markov chain means for source strengths, x-coordinates and y-coordinates. For each panel, chain means appear to converge to similar regions of the state space. Figure 6 de-

picts Markov chain distributions of source strengths, as well as its x-and y-coordinates, at the final iteration (excluding the spin-up period) for the relative times points  $t_0 = 0$  s,  $t_1 = 20$  s and  $t_2 = 40$  s. Their shapes, for each parameter-time combination, appear to be roughly the same despite some noise. However, this is not always an evidence of “good” intra-chain mixing because the marginal prior standard deviations defined earlier for  $\mathbf{b}$  and  $\tilde{\mathbf{a}}$  appear to be “small” compared with the width and height of the aquarium under consideration. For a particular parameter-time combination, one could argue that Markov chain distributions mimic each other due to their starting positions being too close together instead of because they all remain invariant when more particles are obtained.

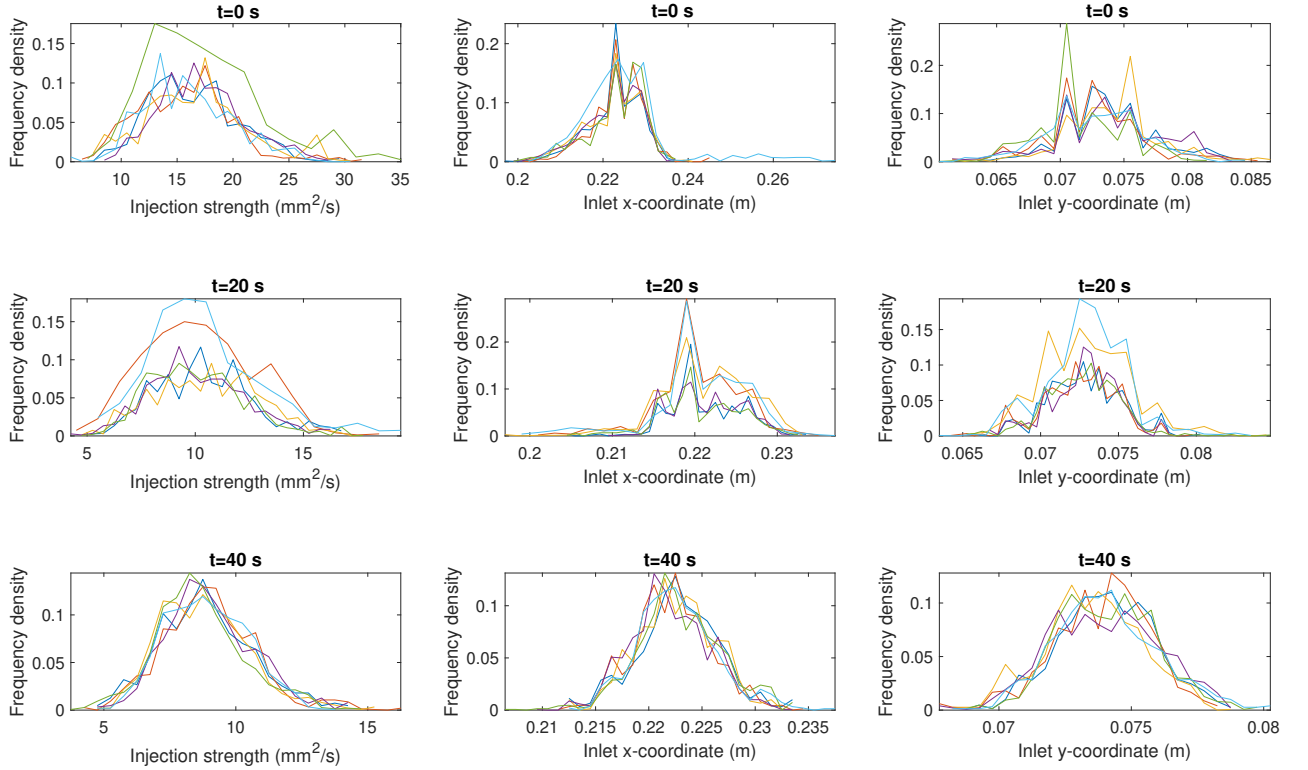


Figure 6: Markov chain distributions at the final iteration for source strengths (first column), x-coordinates (second column) and y-coordinates (third column). The relative times  $t_0 = 0$  s,  $t_1 = 20$  s and  $t_2 = 40$  s constitute the first, second and third rows (respectively). Note that the particles constituting the spin-up period are ignored.

The within-and between-chains variances are now examined. The evolving  $\hat{\mathbf{R}}$  ratios are presented in Figure 7 at the relative times  $t_0 = 0$  s,  $t_1 = 20$  s and  $t_2 = 40$  s. These appear to satisfy the condition  $\hat{\mathbf{R}} < 1.1$  as suggested in [23] upon reaching the final iteration. This, alongside Figures 5 and 6, reasonably suggest that Markov chains have mixed and converged to their relevant marginal posterior distributions.

There appear to be a number of situations in Figures 8 and 9 where within-and between-chains variances stabilise after the end of the spin-up period. There seem to be some exceptions to this as



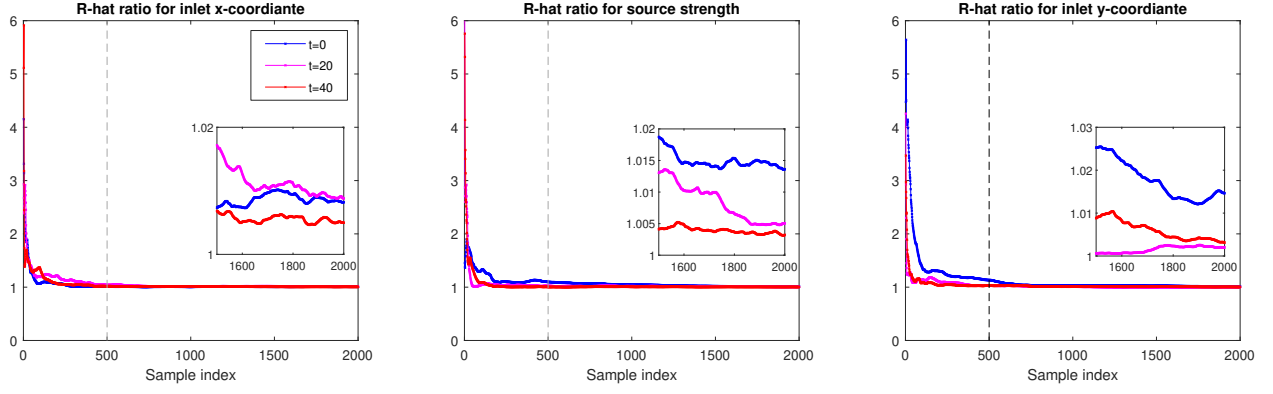


Figure 7: Evolving  $\hat{R}$  ratios for source strengths (first column), x-coordinates (second column), and y-coordinates (third column) at the relative times  $t_0 = 0$  s,  $t_1 = 20$  s and  $t_2 = 40$  s. The vertical dashed line denotes the end of the spin-up period. The horizontal dashed lines denote the threshold  $\hat{R} < 1.1$  suggested in [23].

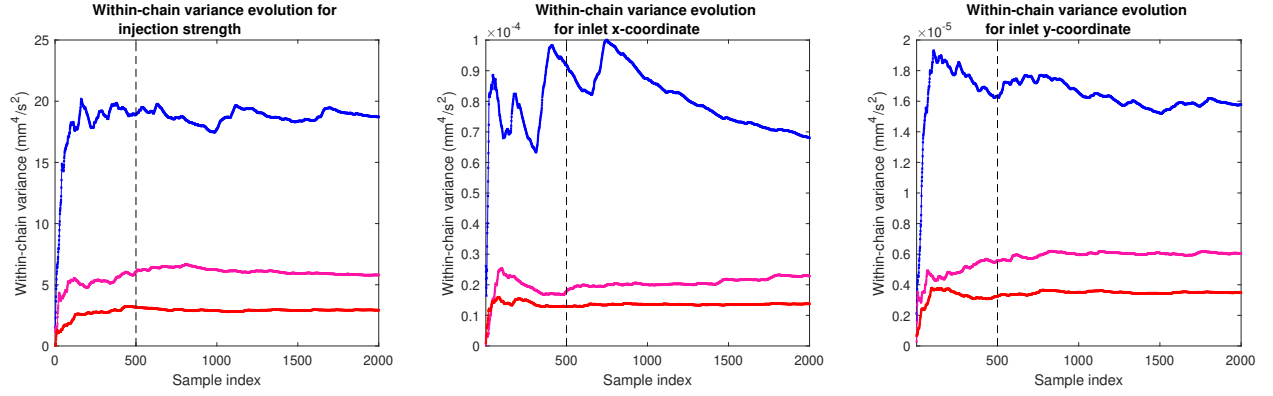


Figure 8: Evolving within-chains variances in the log-log scale for source strengths (first column), x-coordinates (second column), and y-coordinates (third column) at the relative times  $t_0 = 0$  s,  $t_1 = 20$  s, and  $t_2 = 40$  s. The dashed line denotes the end of the spin-up period.

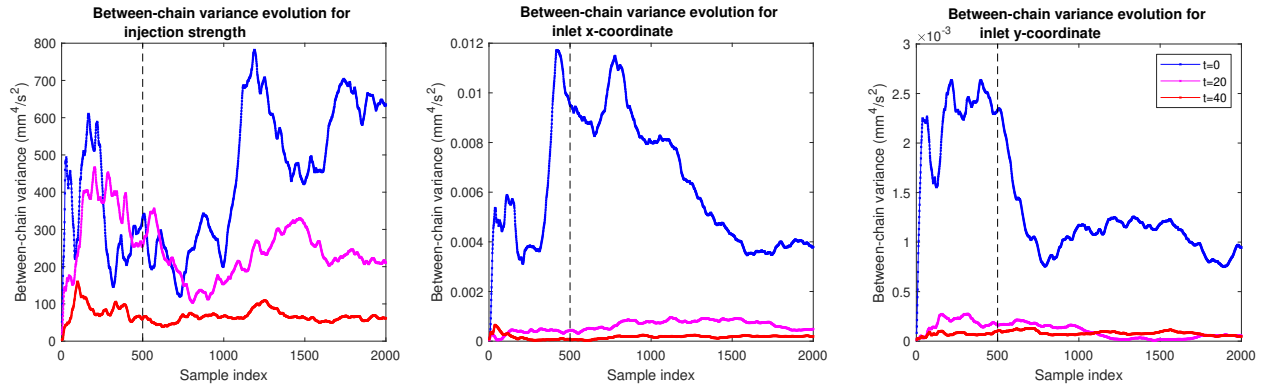


Figure 9: Evolving between-chains variances in the log-log scale for source strengths (first column), x-coordinates (second column), and y-coordinates (third column) at the relative times  $t_0 = 0$  s,  $t_1 = 20$  s, and  $t_2 = 40$  s. The dashed line denotes the end of the spin-up period.

well. For example, the between-chains variances of source strengths at the relative time  $t_2 = 40$  s rapidly increase at later iterations. One possible explanation for this behaviour in the bottom left-hand panel of Figure 9 is that the posterior variance reduction at earlier times could potentially produce “small” marginal prior variances. Starting positions of Markov chains could end up being

so “close” together that their evolving means begin to diverge slightly after the spin-up period has elapsed.

To examine temporal dynamics of optical flows, the flow speed at every grid cell is computed. The frequency distributions of flow speeds at each relative time point in Figure 10 are then constructed and compared with those computed by the multi-resolution Horn-Schunk. These flow speeds are consistent with previous research [22, 48].

The results from the HS and our MCMC approach differ substantially, with a temporally more variable and slower flow being computed using HS (Figure 10). Active bioirrigation by *Arenicola marina* typically occurs over periods of 15 min or longer (e.g. [46, 52]). This suggests that the estimates of our MCMC approach are superior to those obtained using HS, where pumping is estimated to vary substantially over a period of less than a minute. Yet porewater pressure time series measurements reveal variations at high frequencies during periods of pumping, and have revealed strong variations associated with burrowing and backward pumping [48]. Thus, while less common and hence less likely to be captured in our image data, variations in flow magnitudes on timescales shorter than minutes as indicated by the HS approach cannot be ruled out.

Contrary to what is expected based on our understanding of the biology of *Arenicola marina* which injects fluid at its dead-end burrow, the HS approach shows fastest flow located further and further away from the location of the burrowing organism over time (Figure 11), largely tracking the gradients in the tracer concentration field. In these distal gradient regions, flow is well constrained by the observational data. The source of flow is the pumping by *Arenicola marina*, situated near the injection site. Such an injection site is reflected in the location of the divergence in the MCMC approach, but not in the widely distributed divergence in the HS approach. Finally, with a localized injection site, flow velocities should decrease away from the source as the fluid is spreading out. The MCMC approach correctly simulates these properties, whereas the HS method does not. Thus, our work highlights the importance of the data analysis method when inferring the flow field. In this example, the comparison with biological knowledge clearly indicates the advantage of our MCMC approach over traditional methods.

## 7 Summary and outlook

This work develops a parametric model to estimate the flow velocity in the context of bioirrigation. A localized divergence-free flow velocity is desired in the vicinity of the injection point, which

Relative Time (s)	$\hat{c}$ ( $\text{m}^2 \text{s}^{-1}$ )	$\hat{a}$ (m)	$\hat{b}$ (m)	$s^c$ ( $\text{m}^2 \text{s}^{-1}$ )	$s^a$ (m)	$s^b$ (m)
0	$1.6574 \times 10^{-5}$	0.2233	0.0731	$4.3387 \times 10^{-6}$	0.0075	0.0038
10	$1.2861 \times 10^{-5}$	0.2210	0.0729	$3.3386 \times 10^{-6}$	0.0062	0.0034
20	$1.0293 \times 10^{-5}$	0.2209	0.0728	$2.3719 \times 10^{-6}$	0.0049	0.0025
30	$9.2769 \times 10^{-6}$	0.2217	0.0732	$1.9131 \times 10^{-6}$	0.0044	0.0022
40	$8.8754 \times 10^{-6}$	0.2225	0.0739	$1.6880 \times 10^{-6}$	0.0037	0.0019

Table 1: Markov chain means and standard deviations with respect to both particles and chains (excluding the spin-up period) for source strengths, and its x- and y-coordinates. These are obtained using the Metropolis-Hastings sampler and the first five pairs of consecutive down-sampled and cropped real grey scale images.

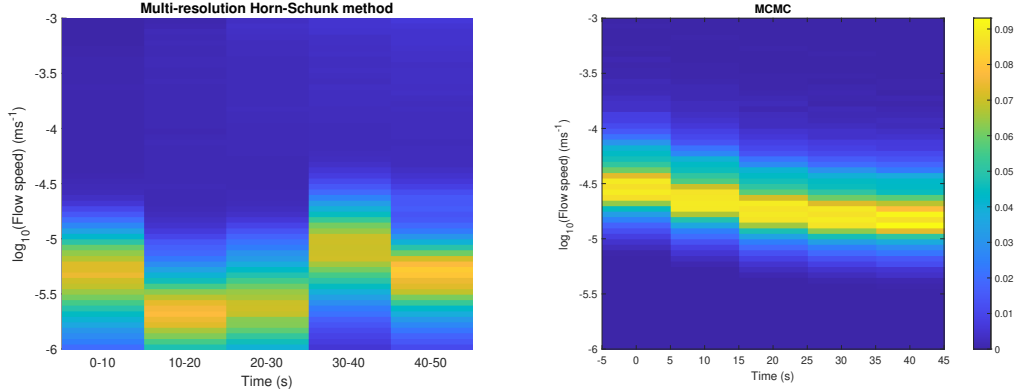


Figure 10: (Left) The frequency distributions of optical flow speeds at the first 5 image pairs computed using the multi-resolution Horn-Schunk, see again Figure 3. (Right) The frequency distributions of optical flow speeds computed using the finite approximation of the ansatz in (5.5), the relation  $\mathbf{u} = \nabla \Phi$  and the parameter estimates in Table 1.

is difficult to achieve with the conventional flow velocity framework method such as the Horn-Schunk method. We demonstrated that the issue of undesired divergence of the flow velocity can be significantly mitigated by using the proposed model in combination with MCMC. The efficacy of the MCMC method in this work indicates the superiority compared to other methods when applied to general problems associated with localized sources of flow/pressure field generation.

Despite the improved result in this initial work, several challenges remain open to overcome. In particular, the model (3.5) can be inaccurate the region further away from the burrow in which the effect of diffusion is considerable and has to be taken into account in the model. This must be addressed if flow velocity estimation has to be extended to the entire area of the image presented in this work. Taking into account diffusion, or considering reactive tracers that inform our understanding of the biogeochemical turnover and efflux of, for example, greenhouse gases such as  $\text{CO}_2$ ,  $\text{N}_2\text{O}$  or  $\text{CH}_4$  from the subsurface would add more parameters, rendering MCMC infeasible in practice. This should motivate a more advanced (approximate) inference to deal with this inverse problem, e.g., iterative Ensemble Kalman filter (IEnKF) for solving inverse problems [8, 15, 38]. Another direction to further improve the estimate by setting up a model to allow a temporal-smoothing

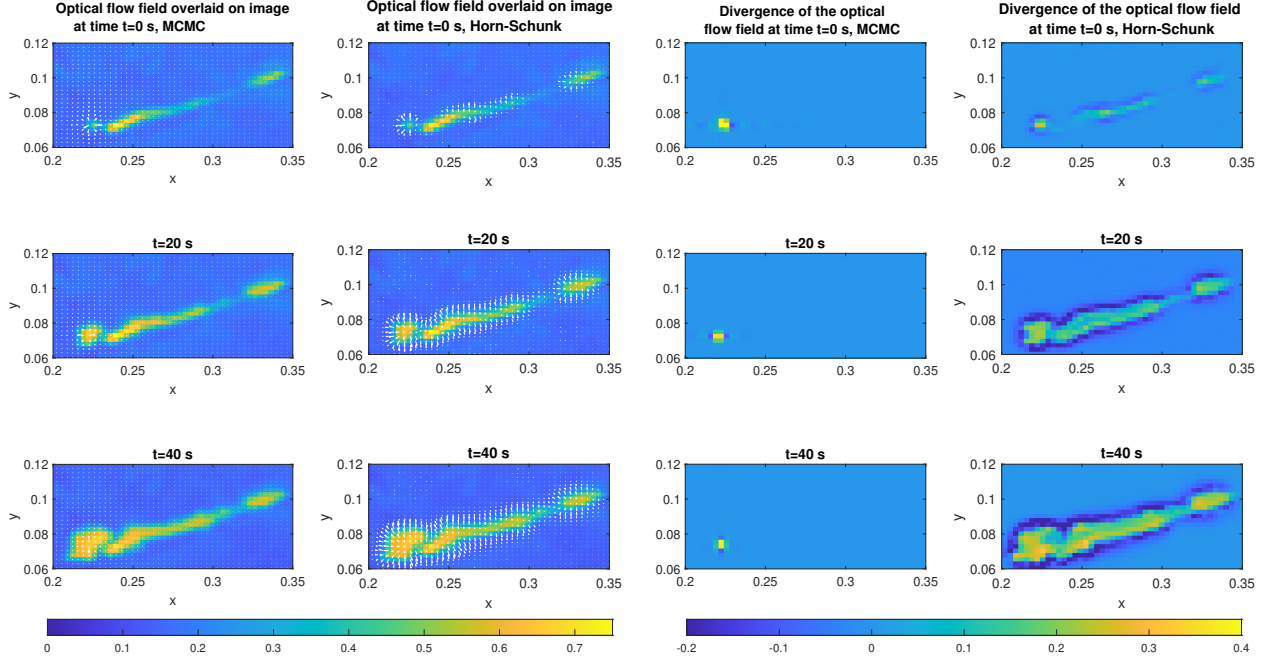


Figure 11: Comparison of the optical flow and divergence between using the finite approximation of the ansatz in (5.5) and Horn-Schunk method in Section 4. The plot shows only the cropped area where *Arenicola marina* is active. The white arrows in the first and second columns denote optical flows. From top to bottom row, the results are shown at the relative times  $t_0 = 0$  s,  $t_1 = 20$  s and  $t_2 = 40$  s, respectively. They are overlaid on top of the down-sampled and cropped real grey scale images associated with those relative times. The third and fourth columns depict the associated divergences. Note that the first and second columns share the same colour scale and likewise for the third and fourth columns.

parameter into the framework instead of using the posterior distribution as a prior distribution in a subsequent image pair. This will in effect avoid “cutting the feedback” issue discussed in [26].

## Acknowledgement

Erik Bollt received funding from the Army Research Office (N68164-EG) and also DARPA. Christof Meile and George Waldbusser were supported by the National Science Foundation (OCE 0751882 to C.M. and OCE 1014226 to G.G.W.). We also thanks to Bing Cheng (University of Surrey) for a fruitful discussion of the derivation of (5.5).

## Appendix A Multi-Resolution Horn-Schunk

To illustrate this method, reconsider the grey scale images  $F_0$  and  $F_1$ . For each image, a pyramid with  $\mathcal{L}$  levels is constructed. At the topmost level (called level 1),  $F_0^1$  and  $F_1^1$  are the “original” grey scale images. Resolution is then reduced by a factor of two with each lower level. Suppose that the prior guess of the optical flow at the bottom pyramid level (i.e.  $\mathcal{L}$ -th level) is zero. This is updated first using (4.6). A more refined optical flow at the  $(\mathcal{L} - 1)$ -th level  $\mathbf{u}^{\mathcal{L}-1}$  is given by

$$\mathbf{u}^{\mathcal{L}-1} = \tilde{\mathbf{u}}^{\mathcal{L}} + \delta \mathbf{u}^{\mathcal{L}-1}, \quad (\text{A.1})$$

where  $\delta \mathbf{u}^{\mathcal{L}-1}(\mathbf{x}, \mathbf{y}, \mathbf{t})$  is a correction increment and  $\tilde{\mathbf{u}}^{\mathcal{L}}(\mathbf{x}, \mathbf{y}, \mathbf{t})$  is obtained by up-sampling  $\mathbf{u}^{\mathcal{L}}$ . A correction increment is obtained by considering the following optimisation problem.

$$\delta \mathbf{u}^{\mathcal{L}-1} = \underset{\delta \mathbf{u}^{\mathcal{L}-1}}{\operatorname{argmin}} \int_{\Omega} \underbrace{(\delta \mathbf{u}^{\mathcal{L}-1} \cdot \nabla F + F_t)^2}_{\text{data term}} + \underbrace{\lambda^2 (\|\nabla(\tilde{\mathbf{u}}^{\mathcal{L}} + \delta \mathbf{u}^{\mathcal{L}-1})\|_2^2 + \|\nabla(\tilde{\mathbf{v}}^{\mathcal{L}} + \delta \mathbf{v}^{\mathcal{L}-1})\|_2^2)}_{\text{regularisation term}} d\Omega. \quad (\text{A.2})$$

Although (A.2) resembles (4.3), there are some differences that require further explanation. Firstly, the data term in (A.2) only involves  $\delta \mathbf{u}^{\mathcal{L}-1}$  because  $\tilde{\mathbf{u}}^{\mathcal{L}}$  ought to already temporally conserve grey scale intensity [21]. Secondly, spatially smooth functions are not closed under addition (depending on their mean). As such, the regularisation term in (A.2) involves both  $\tilde{\mathbf{u}}^{\mathcal{L}}$  and  $\delta \mathbf{u}^{\mathcal{L}-1}$ . Finally, the intensity derivatives in (4.3) are derived from the grey scale images  $F_0$  and  $F_1$  while in (A.2), they are derived from  $F_0^{\mathcal{L}-1}$  and  $F_1^{\mathcal{L}-1}(x - \tilde{u}^{\mathcal{L}} \delta t, y - \tilde{v}^{\mathcal{L}} \delta t, t_1)$ , which can be thought of a warped image of  $F_1^{\mathcal{L}-1}$  towards  $F_0^{\mathcal{L}-1}$  under the action of  $\tilde{\mathbf{u}}^{\mathcal{L}}$ .

Following a derivation analogous to that of (4.6) yields the following iterative Jacobi update scheme:

$$\begin{aligned} \delta u_{i+1}^{\mathcal{L}-1} &= \overline{\overline{\delta u_i}}^{\mathcal{L}-1} - \frac{F_x \left( \overline{\overline{\delta u_i}}^{\mathcal{L}-1} + F_y \overline{\overline{\delta v_i}}^{\mathcal{L}-1} + F_t + F_y \Delta \tilde{v}^{\mathcal{L}} \right)}{\lambda^2 + F_x^2 + F_y^2} + \frac{\lambda^2 + F_y^2}{\lambda^2 + F_x^2 + F_y^2} \Delta \tilde{u}^{\mathcal{L}}. \\ \delta v_{i+1}^{\mathcal{L}-1} &= \overline{\overline{\delta v_i}}^{\mathcal{L}-1} - \frac{F_y \left( \overline{\overline{\delta u_i}}^{\mathcal{L}-1} + F_y \overline{\overline{\delta v_i}}^{\mathcal{L}-1} + F_t + F_x \Delta \tilde{u}^{\mathcal{L}} \right)}{\lambda^2 + F_x^2 + F_y^2} + \frac{\lambda^2 + F_x^2}{\lambda^2 + F_x^2 + F_y^2} \Delta \tilde{v}^{\mathcal{L}}. \end{aligned}$$

We can apply (A.3) to approximate the solution of (A.2). An equivalent form of (A.3), which one

finds easier to work with, is given by the following.

$$\begin{aligned}\delta \mathbf{u}_{i+1}^{\mathcal{L}-1} &= \overline{\overline{\delta \mathbf{u}_i^{\mathcal{L}-1}}} + \Delta \tilde{\mathbf{u}}^{\mathcal{L}} - \frac{\mathbf{F}_x \left( \overline{\overline{\delta \mathbf{u}_i^{\mathcal{L}-1}}} + \mathbf{F}_y \overline{\overline{\delta \mathbf{v}_i^{\mathcal{L}-1}}} + \mathbf{F}_t + \mathbf{F}_y \Delta \tilde{\mathbf{v}}^{\mathcal{L}} \right)}{\lambda^2 + \mathbf{F}_x^2 + \mathbf{F}_y^2} - \frac{\mathbf{F}_x^2}{\lambda^2 + \mathbf{F}_x^2 + \mathbf{F}_y^2} \Delta \tilde{\mathbf{u}}^{\mathcal{L}}. \\ \delta \mathbf{v}_{i+1}^{\mathcal{L}-1} &= \overline{\overline{\delta \mathbf{v}_i^{\mathcal{L}-1}}} + \Delta \tilde{\mathbf{v}}^{\mathcal{L}} - \frac{\mathbf{F}_y \left( \overline{\overline{\delta \mathbf{u}_i^{\mathcal{L}-1}}} + \mathbf{F}_y \overline{\overline{\delta \mathbf{v}_i^{\mathcal{L}-1}}} + \mathbf{F}_t + \mathbf{F}_x \Delta \tilde{\mathbf{u}}^{\mathcal{L}} \right)}{\lambda^2 + \mathbf{F}_x^2 + \mathbf{F}_y^2} - \frac{\mathbf{F}_y^2}{\lambda^2 + \mathbf{F}_x^2 + \mathbf{F}_y^2} \Delta \tilde{\mathbf{v}}^{\mathcal{L}}.\end{aligned}$$

Once  $\tilde{\mathbf{u}}^{\mathcal{L}}$  and  $\delta \mathbf{u}^{\mathcal{L}-1}$  have been recovered, one can then obtain  $\mathbf{u}^{\mathcal{L}-1}$  using (A.1) and up-sample it to level  $\mathcal{L} - 2$ . Then another correction increment can be obtained. This process continues until the topmost level is reached, and flow resolution is the same as that of the “original” images.

Consider another grey scale image  $F_2(x, y, t_2)$ , which is then used in conjunction with  $F_1$  to construct another optical flow. Now suppose that temporal dynamics of optical flows arising from the greyscale images  $\{F_0, F_1, F_2\}$  is unknown, i.e. one cannot write down a model governing their evolution. Then when considering  $F_1$  and  $F_2$ , and intuitive prior guess of the optical flow at the bottom pyramid level would be the most recent estimate which has been down-sampled accordingly. This flow field, or its Laplacian, could be non-zero. If (4.6) was applied at the bottom pyramid level, then this could falsely suggest that the Laplacian of the “incoming” flow (i.e. the prior guess) is zero. As such, at subsequent times, the Horn-Schunck method is not employed at the bottom pyramid level here. The latter of the image pair in question is warped towards the former under the action of the prior guess and (A.3) is applied instead.

## Appendix B Derivation of (5.5)

Solving (5.3) using separation of variables begins with writing the velocity potential being sought in the form

$$\Phi(x, y) = X(x)Y(y). \quad (\text{B.1})$$

The pressure boundary conditions are depicted in Figure 12.

Based on these boundary conditions, the standard separation of variable gives the solution of the form:

$$\Phi(x, y) = \sum_{j=0}^{\infty} \sum_{k=0}^{\infty} \theta^{jk} \cos\left(\frac{k\pi}{L_x} x\right) \cos\left(\frac{(1+2j)\pi}{2L_y} y\right). \quad (\text{B.2})$$

To obtain  $\theta^{jk}$ , we substitute (B.2) to (5.2), revealing



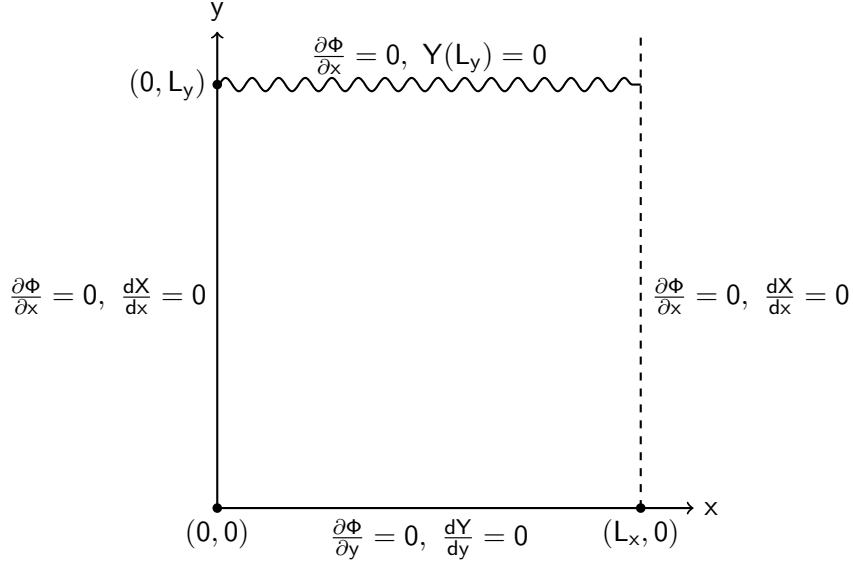


Figure 12: A schematic depicting the boundary conditions in terms of a velocity potential. Dynamics pertaining to the point of injection are neglected.

$$c\delta(x - a, y - b) = \sum_{j=0}^{\infty} \sum_{k=0}^{\infty} \theta^{jk} \left[ -\left(\frac{k\pi}{L_x}\right)^2 - \left(\frac{(1+2j)\pi}{2L_y}\right)^2 \right] \cos\left(\frac{k\pi}{L_x} x\right) \cos\left(\frac{(1+2j)\pi}{2L_y} y\right). \quad (\text{B.3})$$

The above equation can be solved to obtain  $\theta^{jk}$ :

$$\begin{aligned} \theta^{jk} &= \frac{4c}{L_x L_y \nu^{jk}} \cos\left(\frac{k\pi}{L_x} a\right) \cos\left(\frac{(1+2j)\pi}{2L_y} b\right) \\ &= -\frac{4c}{L_x L_y \left[ \left(\frac{k\pi}{L_x}\right)^2 + \left(\frac{(1+2j)\pi}{2L_y}\right)^2 \right]} \cos\left(\frac{k\pi}{L_x} a\right) \cos\left(\frac{(1+2j)\pi}{2L_y} b\right), \end{aligned} \quad (\text{B.4})$$

This expression for  $\theta^{jk}$  is substituted back into (B.2), revealing (5.5).

## References

- [1] R. Basnayake. *Inverse Problems for Image Processing of Spatiotemporal Dynamical Systems*. PhD thesis, Clarkson University, 2014.
- [2] G. Batchelor. *An Introduction to Fluid Dynamics*. Cambridge Mathematical Library. Cambridge University Press, 2000.
- [3] S. Beauchemin and J. Barron. The computation of optical flow. *ACM computing surveys (CSUR)*, 27(3):433–466, 1995.

- [4] K. Bemis, R. Lowell, and A. Farough. Diffuse flow on and around hydrothermal vents at mid-ocean ridges. *Oceanography*, 25(1):182–191, 2012.
- [5] M. Black and P. Anandan. The Robust Estimation of Multiple Motions: Parametric and Piecewise-Smooth Flow Fields. *Computer Vision and Image Understanding*, 63(1):75–104, 1996.
- [6] A. Boetius and F. Wenzhoefer. Seafloor oxygen consumption fuelled by methane from cold seeps. *Nature Geoscience*, 6:725–734, 2013.
- [7] B. Bourdreau. Mean mixed depth of sediments: The wherefore and the why. *Limnology and Oceanography*, 43:524–526, 1998.
- [8] N. Chada, M. Iglesias, L. Roininen, and A. Stuart. Parameterizations for ensemble Kalman inversion. *Inverse Problems*, 34(5):055009, 2018.
- [9] G. Chantas, T. Gkamas, and C. Nikou. Variational-Bayes Optical Flow. *Journal of Mathematical Imaging and Vision*, 50(3):199–213, 2014.
- [10] I. Cohen and I. Herlin. Optical flow and phase portrait methods for environmental satellite image sequences. In *European Conference on Computer Vision*, volume 1065, pages 141–150. Springer, 1996.
- [11] I. Cohen and I. Herling. Nonuniform multiresolution method for optical flow and phase portrait models: Environmental applications. *International Journal of Computer Vision*, 33(1):29–49, 1999.
- [12] T. Corpetti, D. Heitz, G. Arroyo, E. Mémin, and A. Santa-Cruz. Fluid experimental flow estimation based on an optical-flow scheme. *Experiments in Fluids*, 40(1):80–97, 2006.
- [13] C. Cullbertson, S. Jacobson, and R. JM. Diffusion coefficient measurements in microfluidic devices. *Talanta*, 56:365–373, 2002.
- [14] K. Du Clos. Visualizing subsurface burrowing by the polychaete *Alitta virens* with particle image velocimetry. *Limnology and Oceanography: Methods*, 12(10):703–712, 2014.
- [15] A. Emerick and A. Reynolds. Investigation of the sampling performance of ensemble-based methods with a simple reservoir model. *Computational Geosciences*, 17(2):325–350, 2013.
- [16] S. Forster, R. Glud, J. Gundersen, and M. Huettel. In situ study of bromide tracer and oxygen flux in coastal sediments. *Estuarine, Coastal and Shelf Science*, 49(6):813 – 827, 1999.

- [17] D. Fortun, P. Bouthemy, and C. Kervrann. Optical flow modeling and computation: A survey. *Computer Vision and Image Understanding*, 134:1–21, 2015.
- [18] A. Gelman and D. Rubin. Inference from Iterative Simulation Using Multiple Sequences. *Statistical Science*, 7(4):457–472, 1992.
- [19] W. Hastings. Monte Carlo sampling methods using Markov chains and their applications. *Biometrika*, 57(1):97–109, 1970.
- [20] D. Heitz, E. Mémin, and C. Schnörr. Variational fluid flow measurements from image sequences: synopsis and perspectives. *Experiments in Fluids*, 48(3):369–393, 2010.
- [21] B. Horn and B. Schunck. Determining optical flow. *Artificial Intelligence*, 17(1-3):185–203, 1981.
- [22] S. Kaza. Quantification of bioirrigation using image analysis. Master’s thesis, UGA, 2012.
- [23] B. Lambert. *A Student’s Guide to Bayesian Statistics*. Sage, 2018.
- [24] B. Lucas and T. Kanade. An Iterative Image Registration Technique with an Application to Stereo Vision. In *7th Intl Joint Conference on Artificial Intelligence*, pages 121–130, 1981.
- [25] A. Luttmann, E. Bollt, R. Basnayake, S. Kramer, and N. Tuffiaro. A framework for estimating potential fluid flow from digital imagery. *Chaos: An Interdisciplinary Journal of Nonlinear Science*, 23(3), 2013.
- [26] P. M. Cuts in bayesian graphical models. *Statistics and Computing*, 25:37–43, 2015.
- [27] J. Marcello, E. F., and M. F. Methodology for the estimation of ocean surface currents using region matching and differential algorithms. In *IEEE Interantional Geoscience and Remote Sensing Symposium*, pages 882–885, 2007.
- [28] B. McCane, K. Novins, D. Crannitch, and B. Galvin. On Benchmarking Optical Flow. *Computer Vision and Image Understanding*, 84(1):126–143, 2001.
- [29] D. L. McCormack. *Parameter Estimation and Inverse Problems for Reactive Transport Models in Bioirrigated Sediments*. PhD thesis, University of Surrey, 2019.
- [30] C. Meile and P. Cappellen. Global estimates of enhanced solute transport in marine sediments. *Limnology and Oceanography*, 48(2):777–786, 2003.

- [31] C. Meile, C. Koretsky, and P. Cappellen. Quantifying bioirrigation in aquatic sediments: An inverse modeling approach. *Limnology and Oceanography*, 46(1):164–177, 2001.
- [32] F. Meysman, O. Galaktionov, B. Gribsholt, and J. Middelburg. Bio-irrigation in permeable sediments: An assessment of model complexity. *Journal of Marine Research*, 64(4):589–627, 2006.
- [33] F. Meysman, O. Galaktionov, B. Gribsholt, and J. Middelburg. Bioirrigation in permeable sediments: Advective pore-water transport induced by burrow ventilation. *Limnology and Oceanography*, 51(1):142–156, 2006.
- [34] L. Polerecky, N. Volkerborn, and P. Stief. High temporal resolution oxygen imaging in bioirrigated sediments. *Environmental Science & Technology*, 40(18):5763–5769, 2006.
- [35] H. Riisgård and G. Banta. Irrigation and deposit feeding by the lugworm *Arenicola marina*, characteristics and secondary effects on the environment. A review of current knowledge. *Vie Milieu*, 48(4):243–257, 1998.
- [36] H. Riisgård, I. Berntsen, and B. Tarp. The lugworm (*Arenicola marina*) pump: characteristics, modelling and energy cost. *Marine Ecology Progress Series*, 138:149–156, 1996.
- [37] P. Ruhnau, T. Kohlberger, C. Schnörr, and H. Nobach. Variational optical flow estimation for particle image velocimetry. *Experiments in Fluids*, 38(1):21–32, 2005.
- [38] P. Sakov, J. Haussaire, and M. Bocquet. An iterative ensemble Kalman filter in the presence of additive model error. *Quarterly Journal of the Royal Meteorological Society*, 144(713):1297–1309, 2017.
- [39] N. Santitissadeekorn and E. Bollt. The infinitesimal operator for the semigroup of the Frobenius-Perron operator from image sequence data: Vector fields and transport barriers from movies. *Chaos: An Interdisciplinary Journal of Nonlinear Science*, 17(2):023126, 2007.
- [40] I. Santos, B. Eyre, and M. Huettel. The driving forces of porewater and groundwater flow in permeable coastal sediments: A review. *Estuarine, Coastal and Shelf Science*, 98:1–15, 2012.
- [41] K. Shum. Wave-induced advective transport below a rippled water-sediment interface. *Journal of Geophysical Research*, 97(C1):789–808, 1992.

- [42] E. Simoncelli and E. Adelson. *Computing Optical Flow Distributions Using Spatio-temporal Filters*. Vision and Modeling Group, Media Laboratory, Massachusetts Institute of Technology, 1991.
- [43] E. Simoncelli and E. Adelson. *Computing Optical Flow Distributions Using Spatio-temporal Filters*. Citeseer, 1991.
- [44] M. Solan, E. Ward, E. White, and et al. Worldwide measurements of bioturbation intensity, ventilation rate, and the mixing depth of marine sediments. *Scientific Data*, 6(58), 2019.
- [45] J. Sun, J. Fernando, and E. Bollt. Bayesian optical flow with uncertainty quantification. *Inverse Problems*, 34(10):105008, 2018.
- [46] K. Timmerman, G. Banta, and G. RN. Linking arenicola marina irrigation behavior to oxygen transport and dynamics in sandy sediments. *Journal of Marine Research*, 64(6):915–938, 2006.
- [47] N. Volkenborn, S. Hedtkamp, J. van Beusekom, and K. Reise. Effects of bioturbation and bioirrigation by lugworms (*Arenicola marina*) on physical and chemical sediment properties and implications for intertidal habitat succession. *Estuarine, Coastal and Shelf Science*, 74(1-2):331–343, 2007.
- [48] N. Volkenborn, L. Polerecky, D. Wetthey, and S. Woodin. Oscillatory porewater bioadvection in marine sediments induced by hydraulic activities of *Arenicola marina*. *Limnology and Oceanography*, 55(3):1231–1247, 2010.
- [49] N. Volkenborn, S. Woodin, D. Wetthey, and L. Polereck. *Reference Module in Earth Systems and Environmental Sciences*, chapter Bioirrigation in Marine Sediments. Elsevier, 2016.
- [50] G. Waldbusser and R. Marinelli. Macrofaunal modification of porewater advection: role of species function, species interaction, and kinetics. *Marine Ecology Progress Series*, 311:217–231, 2006.
- [51] K. William. *Digital image processing*. Wiley-Interscience, 4 edition, 2007.
- [52] S. Wohlgemuth, A. Taylor, and M. Grieshaber. Ventilatory and metabolic responses to hypoxia and sulphide in the lugworm arenicola marina (l.). *The Journal of Experimental Biology*, 203:3177–3188, 2002.
- [53] S. A. Woodin and W. D. S. Arenicolid behaviors: similarity of arenicola marina and abarenicola pacifica. *Zoosymposia*, 2:447–456, 2009.

- [54] S. A. Woodin, D. S. Wethey, and N. Volkenborn. Infernal hydraulic ecosystem engineers: cast of characters and impacts. *Integrative and Comparative Biology*, pages 1–12, 2009.